

**RELATIONSHIP BETWEEN IRON VALENCE STATES OF SERPENTINE IN CM CHONDRITES AND THEIR AQUEOUS ALTERATION DEGREES.** T. Mikouchi<sup>1</sup>, M. Zolensky<sup>2</sup>, W. Satake<sup>1</sup>, and L. Le<sup>3</sup>, <sup>1</sup>Dept. of Earth and Planetary Science, University of Tokyo, Hongo, Tokyo 113-0033, Japan (mikouchi@eps.s.u-tokyo.ac.jp), <sup>2</sup>Astromaterials Research and Exploration Science, KT, NASA Johnson Space Center, Houston, TX 77058, USA, <sup>3</sup>Jacobs ESCG, Houston, TX 77058, USA.

The 0.6-0.7  $\mu\text{m}$  absorption band observed for C-type asteroids is caused by the presence of  $\text{Fe}^{3+}$  in phyllosilicates [1]. Because Fe-bearing phyllosilicates, especially serpentine, are the most dominant product of aqueous alteration in the most abundant carbonaceous chondrites, CM chondrites [e.g., 2,3], it is important to understand the crystal chemistry of serpentine in CM chondrites to better understand spectral features of C-type asteroids. CM chondrites show variable degrees of aqueous alteration [4,5], which should be related to iron valences in serpentine. It is predicted that the  $\text{Fe}^{3+}/\Sigma\text{Fe}$  ratios of serpentine in CM chondrites decrease as alteration proceeds by Si and  $\text{Fe}^{3+}$  substitutions from end-member cronstedtite to serpentine [4], which should be apparent in the absorption intensity of the 0.6-0.7  $\mu\text{m}$  band from C-type asteroids. In fact, the JAXA Hayabusa 2 target (C-type asteroid: 1993 JU3) exhibits heterogeneous spectral features (0.7  $\mu\text{m}$  absorption band disappears by rotation) [6].

From these points of view, we have analyzed iron valences of matrix serpentine in several CM chondrites which span the entire observed range of aqueous alteration using Synchrotron Radiation X-ray Absorption Near-Edge Structure (SR-XANES). In this abstract we discuss the relationship between obtained  $\text{Fe}^{3+}/\Sigma\text{Fe}$  ratios and alteration degrees by adding new data to our previous studies [7,8].

We have so far analyzed Murray, Nogoya, ALH 84029 [7], Murchison, Cold Bokkeveld and a clast of CM1 lithology within Tagish Lake (thin section KN1) [8], and we newly analyzed Kivesvaara (Table 1). The SR-XANES analyses were performed at BL-4A of the Photon Factory (PF), KEK in Tsukuba, Japan. The beam size was *ca.* 6 x 5  $\mu\text{m}$ . We used kaersutite amphibole for the  $\text{Fe}^{2+}$  and  $\text{Fe}^{3+}$  standards whose  $\text{Fe}^{3+}/\Sigma\text{Fe}$  ratios were determined by wet chemistry [9], and estimated the  $\text{Fe}^{3+}/\Sigma\text{Fe}$  ratio of CM serpentine by a linear relationship between the centroid energy position of XANES Fe K pre-edge spectra and the  $\text{Fe}^{3+}/\Sigma\text{Fe}$  ratio ( $\pm 10\%$  error).

In our previous study, we did not observe clear difference between  $\text{Fe}^{3+}/\Sigma\text{Fe}$  ratios of serpentine and alteration degree, although we analyzed samples showing a wide range of aqueous alteration [7,8]. We saw only limited ranges of  $\text{Fe}^{3+}/\Sigma\text{Fe}$  ratios of serpentine, showing mostly  $\text{Fe}^{3+}$ -rich compositions (Table 1). Our new-

ly-obtained  $\text{Fe}^{3+}/\Sigma\text{Fe}$  ratios of serpentine in Kivesvaara is  $\sim 0.9$ -1 for intermediate Mg-Fe serpentine, but Mg-rich serpentine clearly has a lower  $\text{Fe}^{3+}/\Sigma\text{Fe}$  ratio of  $\sim 0.5$ . Because Kivesvaara is a minimally altered sample among CM chondrites studied, we expect that its  $\text{Fe}^{3+}/\Sigma\text{Fe}$  ratio in serpentine is the most  $\text{Fe}^{3+}$ -rich. However, we again found no correlation between  $\text{Fe}^{3+}/\Sigma\text{Fe}$  ratios of serpentine and alteration degree [7,8].

We suggest that the analyzed serpentine contains submicron Fe oxide or oxyhydroxide phases that affect XANES spectra in some samples. For example, the original serpentine compositions in heavily-weathered samples were heterogeneous ( $\text{Fe}^{2+}$ -rich), but terrestrial oxidation has made much of the  $\text{Fe}^{2+}$  into  $\text{Fe}^{3+}$  to form nano-phase ferrihydrite [3]. In contrast, minimally weathered samples contain small amounts of Mg-Fe anhydrous silicates ( $\text{Fe}^{2+}$ -rich) with  $\text{Fe}^{3+}$ -rich serpentine. Probably, smaller spatial resolution may be required for iron valence analysis of CM serpentine [10].

Table 1. Alteration degrees and  $\text{Fe}^{3+}/\Sigma\text{Fe}$  ratios of serpentine in CM chondrites using SR-XANES.

Sample	Alteration index <sup>[4]</sup>	Petrographic type <sup>[5]</sup>	$\text{Fe}^{3+}/\Sigma\text{Fe}$ (Mg-rich)	$\text{Fe}^{3+}/\Sigma\text{Fe}$ (Intermediate)	$\text{Fe}^{3+}/\Sigma\text{Fe}$ (Fe-rich)
Kivesvaara	-	2.5	0.5	0.90-1.0	-
Murchison <sup>[7]</sup>	1.57	2.5	0.82-0.84	-	0.81
Murray <sup>[6]</sup>	1.43	2.4/2.5	0.87-0.88	-	0.76-0.88
Nogoya <sup>[6]</sup>	1.03	2.2	0.82-0.84	0.85-0.90	0.75-0.79
Cold Bokkeveld <sup>[7]</sup>	0.97	2.2	0.80	0.83	0.84
ALH84029 <sup>[6]</sup>	-	-	0.87-0.94	-	0.18-0.73
Tagish Lake (KN1) <sup>[7]</sup>	-	1	0.80-0.82	-	0.78

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